

## RESISTIVITY AND VLF MEASUREMENTS FOR SPRING MECHANISM DETERMINATION AT NE CHIOS ISL.

<sup>1</sup>PAPADOPOULOS T.D., <sup>1</sup>ALEXOPOULOS J.D., <sup>1</sup>DILALOS S., <sup>2</sup>PIPPIDIS M.J.

<sup>1</sup>NKUA, Faculty of Geology and Geoenvironment, Dep. Geophysics-Geothermics

<sup>2</sup>Environmental Agency of Xios Prefecture

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### Abstract

In this paper the feeding mechanism of springs distributed around the broader area of Delfini at NE side of Chios island is explored, in an effort to locate possible fresh water paths inland before they are discharged at the sea. The drilling of more production hydro wells is of vital importance in order to cover the drinking and irrigation demands. Although Chios has been extensively drilled, the demand for further water investigations comes from the need of increased water consumption during summer time in one hand and from the inefficiency of alternative water sources (e.g. construction of dams, desalination units, etc) in the other hand. Geophysical methods were used (VLF and resistivity) to explore in detail water potential areas, like the broader area of Delfini, where huge quantities of fresh water discharges to the sea level. The presence of a confined aquifer composed mainly from limestones which are overlain by clastic impermeable deposits, together with the existence of water potential bearing fracture zones in the nearby area determine the hydrogeological regime and explain the positive elevations of brakish waters found during summer time. It seems that the lateral extent of impermeable clastic deposits plays a crucial role in allowing the inland flow of sea water and combined with the presence of the confined aquifer produce an upward mixing flow of fresh and sea water.

### 1. Introduction

The island of Chios has been extensively drilled for water exploitation the last two decades, in an effort to cover the needs of water supply in urban and rural areas. The demand for greater quantities of water consumptions increases particularly during summer. The hydrogeological basin of "Korakari" is considered as the most water productive area in Chios Isl., providing >80% of the water supply and >70% covering of the irrigation needs of the island (Μπάνος *et al.* 1995). Although it seems that the water supply problem is not intense, however some production wells appear an abnormal increase in mercury (Hg) content, around the area of "Korakari" during summer. This problem requires the investigation of other potentially productive areas. Of course, local authorities and governmental agencies have also proposed other more permanent solutions like the construction of dams (e.g. Katraris, Fyrolakas, Koris Gefyri, Sarapio, etc.) and the establishment of desalination units. The construction of dams started long time ago but still are out of operation, mainly because of financial problems. Desalination units were also put forward in the area of Lagada in NE Chios but their maintenance cost is high so they were not used in other places. Besides these intermediate and long term solutions, other approaches were proposed to cover immediate needs and solve the water supply problem in Chios Island.

A geophysical survey was proposed to search for new potentially promising areas for further drilling. These areas must fulfill the following prerequisites: a) To be in regions that are not extensively drilled by previous works, b) to be in accessible places in order to be reached easily by a drilling machine, c) to be in areas that are promising from a hydrogeological point of view based on the detailed consideration of the geological structure and the potentiality of the geophysical methods used and d) to be close to the existing water pipe network for direct connection of the new productive wells. As it is obvious the constraints mentioned above, make the problem more complicated and impose specific geophysical methods to be used.

In order to detect vertical or sub-vertical conductive zones the VLF method was applied. From the areas that were explored the region of Delfini was extensively surveyed. Great surface water discharges appear close to sea level (SP1 at +4.2m elevation, SP2 +2.6 to +0.9 and SP3  $\pm 0$ ; Figure 1) in the broader area of Delfini. The idea was to study possible water paths through karstic structures or fault zones before the water discharges reach the sea level. The detailed geological mapping of the area under investigation helped us to draw the appropriate geophysical profiles. Besides VLF measurements, resistivity soundings were also conducted to investigate sea water intrusion and the depth of carbonate rocks under the alluvial deposits in Delfini plain.

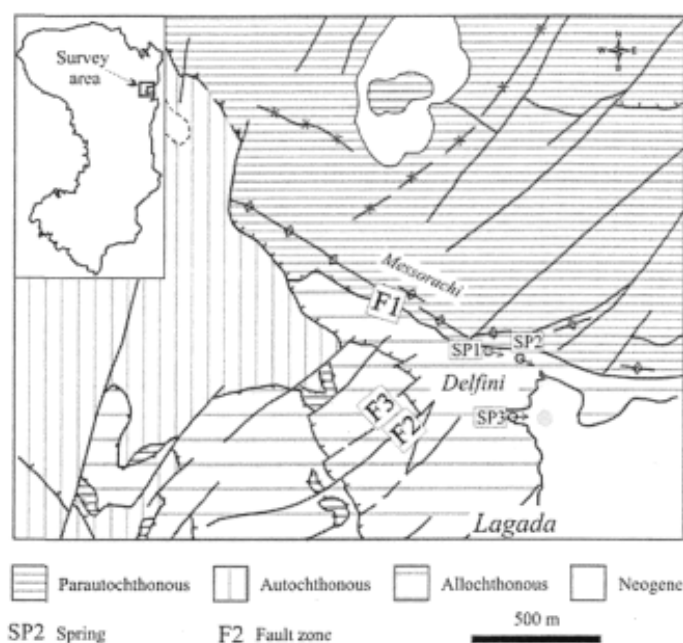


Figure 1. Major tectonostratigraphic units of the survey area (Kauffmann G. 1969).

## 2. Geological setting

Chios Island can be subdivided into two tectonostratigraphic units (Besenecker *et al.* 1968):

First, an autochthonous unit (Figure 1) which consists from the base of mainly clastic rocks of late Palaeozoic age, that are overlain by a succession of early Triassic conglomerates, sandstones followed from well bedded limestones of late early Triassic age and massive limestones interbedded with reddish Hallstatt type limestones. This sequence is overlain by an Anisian volcano-sedimentary succession, followed by the well bedded marly limestones of late Anisian and a more than 1000 meters thick Ladinian to Rhaetian-Liassic age limestones, with few clastic intercalations (*emersion horizons*).

Second, an allochthonous unit (Figure 1) which overlies parts of the autochthonous-parautochthonous basement and consists of a lower horizon of quartzose graywakes, sandstones and minor siltstones of late Carboniferous age, followed by limestones and a clastic carbonaceous succession of early Permian age. The middle Permian is represented by carbonate sequences, which are overlain by transgressive sediments (conglomerates, sandstones, siltstones followed by thick bedded Dachstein-type limestones and minor dolomites) of early Jurassic age. Tectonically between these two units are caught parts of the autochthonous succession as parautochthonous unit. The formations of the Aepos-Pytious karstified plateau which comprise



north. Parallel to it, shorter faults intersect the allochthonous formations of the Delfini depression and the Skoukla peninsula. In NW to SE direction lies the Anavatos–Nea Moni fault zone.

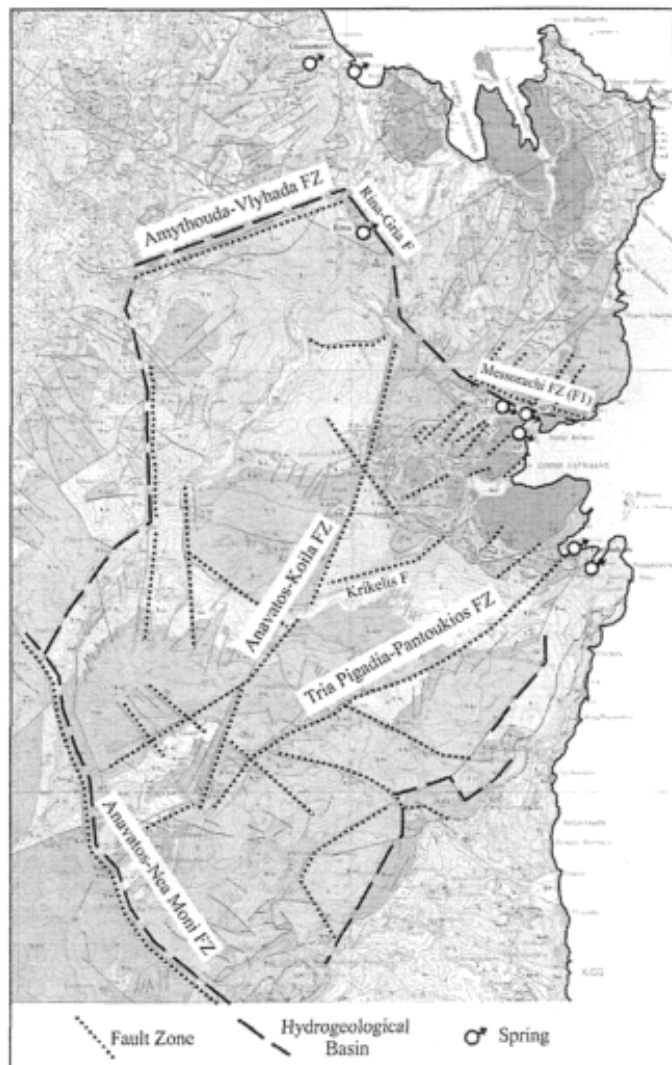


Figure 3. The broader ground water basin of the area under study along with the main fault zones (modified from geological map "Chios", Institute for Geology and Subsurface Research 1971).

### 2.1 Hydrology and watershed of the Delfini karstic coastal system

The central part up to NE of Chios Isl. is covered by a large limestone plateau of Jurassic–Triassic age, with a complicated tectonic structure. The limestone plateau, stays from about >800m WSW to 400m ENE above sea level and its north eastern termination (the Delfini–Kardamyla depression) from >200m to below sea level (the depth of the sea bottom, which forms an underwater plain in front of the depression is constantly 70 to 80 m). The major part of

the plateau is intensely karstified and the annual rainfalls exceed 700 mm/y (>90% of the total rain) during the winter months, from October to April. The majority of the plateau surface is a denudated karst covered of scarce scrub of low vegetation which presents high infiltration rates that are estimated more than 30% and locally >45%. Between the gulfs of Pantoukios and Delfini the watershed dips below sea level, to a depth of about 70-80m. The East-North borders are designated from the contact between the parautochthonous and the allochthonous unit in the Messorachi region, the Delfini tectonic zone and the north part of the Anavatos-Koila fault zone (Figure 3). In the north, the fault zone Amythounda-Vlyhada separates the Delfini hydrogeological unit from the Nagos-Giosonas unit. The watershed basin of the Delfini karstic coastal system, covers more than 90 km<sup>2</sup> and the yearly discharge may exceed 30x10<sup>6</sup> m<sup>3</sup>, a quantity which is almost three times higher than the actual water demand/consumption on the island. The possible borders of the broader ground water basin that feeds with huge water discharges the gulfs of Pantoukios and Delfini are shown in Figure 3.

### 2.2 Hydrogeological aspects of the Delfini aquifer

The Palaeozoic clastic rocks of the autochthonous unit form the impermeable basement. The early Triassic conglomerates, sandstones, the well bedded limestones which follow, the massive limestones which incorporate Hallstatt inclusions and the more than 1000m thick Ladinian to Rhaetian-Liassic limestones, form the potential permeable cover. The Anisian volcano-sedimentary succession and the well bedded marly limestones, together with the emersion horizons can form impermeable layers. The quartzose graywakes, sandstones and siltstones of the lower horizon form a possible discontinuous impermeable layer, together with the clastic carbonaceous succession of the early Permian and the clastic sediments in the base of the thick bedded limestones and dolomites which close the stratigraphic sequence of the allochthonous unit. The Dachstein type thick bedded limestones, of early Jurassic age, form a potentially permeable formation. The numerous intersecting faults favored the development of limestone karstification and produced an efficient subsurface drainage system in which probably the Anavatos-Koila fault zone serves as a main collector, which feeds the Krikelis and the parallel to it faults in the Delfini depression.

### 2.3 Discharge characteristics of the Delfini karstic coastal system

The lower part (also below sea level) of the aquifer in the Delfini depression area which is mainly developed in the autochthonous limestones, is partly (in unknown extent) covered by the impermeable strata of the allochthonous and the parautochthonous formations. This cover protects partly the aquifer which discharges a considerable amount of fresh water at the Giuvari spring (SP3), without mixing with sea water. This spring flows at the intersection of a parallel to Krikelis fault with a thick Quaternary clay cover in the bottom of the small gulf in front of it, at zero elevation. The year round discharge is estimated of more than 0.45x10<sup>6</sup> m<sup>3</sup>. The Pantoukios gulf, which forms the extension of the Tria Pigadia-Pantoukios fault, discharges below sea level (>30m.) big quantities of karst water which are observable at sea surface most of the year. In the Lagada gulf, does not appear any spring over or below sea level. In the north end of the Delfini depression along the Messorachi fault zone area, the Delfini springs (SP1 at +4.2 m elevation and SP2 at different locations from +2,6 to +0,9 m above sea level), discharge more than 5x10<sup>6</sup> m<sup>3</sup>/annually. The fluctuation of the volume and the salinity of the springs depend on the annual rate of the rainfalls, the location of the springs and the elevation of the discharge. Submarine springs along that fault line are not known. On the north end of the watershed, the Rina springs at an elevation of +56m discharge (only during the months of the winter a quantity that can vary considerably from 0 to more than 1x10<sup>6</sup> m<sup>3</sup>/y) at the intersection of a parallel to Anavatos-Koila fault zone with the Rina-Gria fault which is parallel to Messorachi fault zone.

#### 2.4 Investigation of the Delfini discharge zone

Two main springs discharge in the area of Delfini (SP1 and SP2) near the sea level (Figures 2 and 3). Both springs discharge great quantities of fresh water especially during winter and spring times with very good quality of water. The municipality of Chios through the Water Management Organization (DEYAX) has built aqueducts at spring sites SP1 and SP3 (Giuvari) and supplies the water network of the island. During summer time the great quantities of discharge run short and the water quality is downgraded as it is affected by the sea water intrusion. The feeding of both springs was an unresolved problem and needed further investigation. In one hand, it seems that the fault zone F1 (Figures 1 and 2) is responsible for the discharge at site SP1 and on the other hand the parautochthonous series on the NE side of the fault zone does not show water feeding potential since the upper carbonate rocks of the unit are of small extent. Another interesting point is that the spring at site SP1 during summer time (low discharges) presents higher salinity values although lies at positive altitudes (~4.2m). The allochthonous series on the SW side of fault zone F1, is generally considered as a non potentially water exploitation region since it is composed from impermeable rocks.

It seems that the springs SP1, SP2 and SP3 (Figures 2 and 3), share a common feeding line, of Krikelis fault and the subparallel to it of smaller faults (e.g. F2 and F3) which are well developed in the Delfini depression and are in some way connected to Anavatos-Koila fault zone. The Messorachi fault zone (F1), intersects almost perpendicularly those faults and probably brings the sea water inland. The faults must affect the main aquifer on the autochthonous basement which is effectively protected from sea water intrusion, mainly by the cover of the allochthonous unit and probably by the recent clay rich deposits, like the case of spring SP3 (Giuvari). This model of the discharge mechanism must be validated prior to establishment of a viable exploitation scheme of that aquifer in the region around the Delfini valleys and at an elevation of less than 100m. So, it is necessary to answer the follow questions: Are the NE to SW faults in the Delfini depression affecting the autochthonous basement? Which of the mapped faults are active hydrologically and in which way they are affected by the sea water intrusion? How deep are the impermeable covers? Is the aquifer limited in the fault zone or not? Is the Messorachi fault zone the canal for sea water intrusion? At what depth the drilling must end up, before getting in contact with a deeper mixed with sea water aquifer?

#### 3. VLF measurements

VLF measurements were conducted in order to detect vertical to sub-vertical conductive zones or karstic structures in carbonate rocks perpendicular to fault zones as it is shown in Figure 2. The spacing between stations was 5 meters and the VLF lines were kept long enough (a length greater than 200m) to outline the VLF anomalies. Although the relief was not smooth along almost all lines, however the filtered data (Walen 2003, Ogilvy and Lee 1991, Karous and Hjelt 1983) provided reliable results without applying topographic corrections.

In the area of Delfini at least five (5) VLF lines were conducted mainly to detect the NE extension of fault zones F2 and F3. Two (2) VLF lines were also carried out transversely to the fault zone F1 to examine its potential relative to SE-ward discharge of Messorachi Mtn. In Figure 4 the interpreted filtered data for the Delfini area are shown. The fault F2 is extended NE-ward (Figure 4) up to the plain of Delfini. Fault F3 does not show a clear NE extension.

#### 4. Resistivity survey

Resistivity measurements were also conducted in the area of Delfini westwards from spring SP1 (Figure 2). Fourteen (14) soundings (VES) were carried out, using the Schlumberger configuration and further three (3) in situ measurements were conducted. The main purpose of this investigation was to outline the inward extension of seawater front and to characterize the nature of the basement in the NE extension of faults F2 and F3. Although the geoelectrical data were processed by applying the automatic method of Zohdy and Bisdorf (Zohdy 1989) and the commercial software packages IX1D (v3) of Interpex (2006) and WinSev of W-GeoSoft, the final interpretation was based on IX1D package. According to these resistivity results (high

resistivity values, >250 Ohm.m), the basement in the area between the soundings D01, D02, D08 is composed of carbonate rocks at depths greater than 20 meters (Figure 5). The area between soundings D03, D09 and D10 presents lower resistivity values (<60 Ohm.m) up to depths of 40 meters.

The geological formation with such resistivity values is compatible with the presence of clastic rocks of the allochthonous series. At depths greater than 40 meters, highly resistive (>300 Ohm.m) formations of probably carbonate nature are present. The geoelectric sounding D04 showed very low resistivity values (~6 Ohm.m) indicating the inward (inland) intrusion of sea water. The area south of sounding D04 is swampy, not allowing further investigation of this area. The interesting point here is that sounding D04 lies at an elevation of approximately 6 meters above sea level and the low (~6 Ohm.m) resistivity value was detected at positive elevation of about 1m below ground surface. The spring SP1 which is close to sounding D04 shows similar behavior.

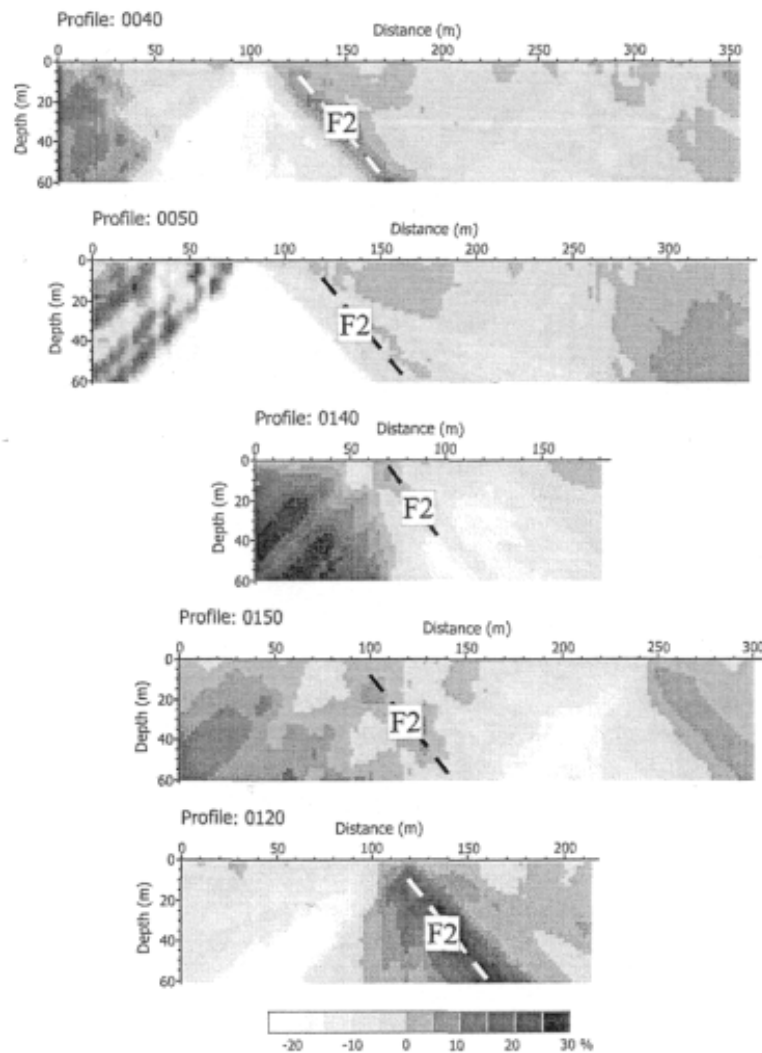


Figure 4. VLF interpreted filtered data in Delfini study area.

All resistivity measurements were carried out during July 2007, at times where fresh water discharges were low. Measurements of water conductivity made in the well W1 (Figure 3) showed high content of chloride ions during August 2007. The well is close to geoelectric sounding D14 (about 10 meters) and its water level was at an elevation of about 5-5.5 meters. The chloride content indicates the affect of seawater intrusion inland and according to resistivity values found in this area, the seawater front must be extended close to the area of geoelectrical soundings D09 and D10.



Figure 5. 3D representation of the geoelectrical basement (limestones / dolomites).

### 5. Integrating geophysical and geological results

In order to integrate geophysical results and geological data obtained from the area of Delfini, two cross sections were drawn in NE-SW (Section 1) and NNE-SSW (Section 2) directions respectively (Figure 2). These two sections are shown in Figure 6. The subsurface geological setting of Delfini plain is composed of five mainly formations, two of which are the shallow clastic deposits (Ci-s) and the deeper clastic rocks (Ci) which are both impermeable ones. The presence of the shallow impermeable clastic deposits keep out the sea water intrusion inland and seems that the only free passage of sea water intrusion is through the alluvial and shore deposits NE-ward of sounding D14. The alluvial deposits in this area extend below sea level, allowing thus the sea water intrusion. Below the alluvial and clastic deposits are extended the permeable limestone-dolomite (Ji) and dolomite-limestone (Tr m) formations that form a confined aquifer. The presence of fault F2 in cross section 2 which is extended further to NE, might discharge great quantities of fresh water in the plain of Delfini. The presence of fault F1 might also discharge quantities of fresh water but if this fault extends further SE to Delfini gulf, it could work as well as a canal for sea water intrusion inland. It is thought that the presence of the impermeable clastic deposits and clastic rocks create a confined aquifer of fresh water affected by saline water especially during summer time when the fresh water discharge is low. This mechanism can explain the presence of brackish waters at positive elevations during summer time.

### 6. Conclusions

The target of drilling new exploitation wells in the broader area of Delfini at NE side of Chios island was fulfilled, by exploring through geophysical methods the hydrogeological regime and detecting sites of fresh water potential. By combining VLF and geoelectrical results along with geological data it was possible to determine the mechanism and lateral feeding of the springs

SP1, SP2 and SP3 which provide huge discharges during winter and spring times. The confined aquifer composed mainly of limestones is overlain by impermeable clastic deposits (Ci-s) which although cover a large part of Delfini plain, allow the fresh water and/or sea water to rise above sea level beyond geoelectrical sounding D14. This is possible since the permeable alluvial deposits lie directly on the limestone basement (Figure 6). The lateral fresh water feeding comes from fault F2 which is extended further to NE in the Delfini plain. Fault F1 might discharge some quantities of fresh water at springs SP1 and SP2 but probably works as a canal for sea water intrusion inland.

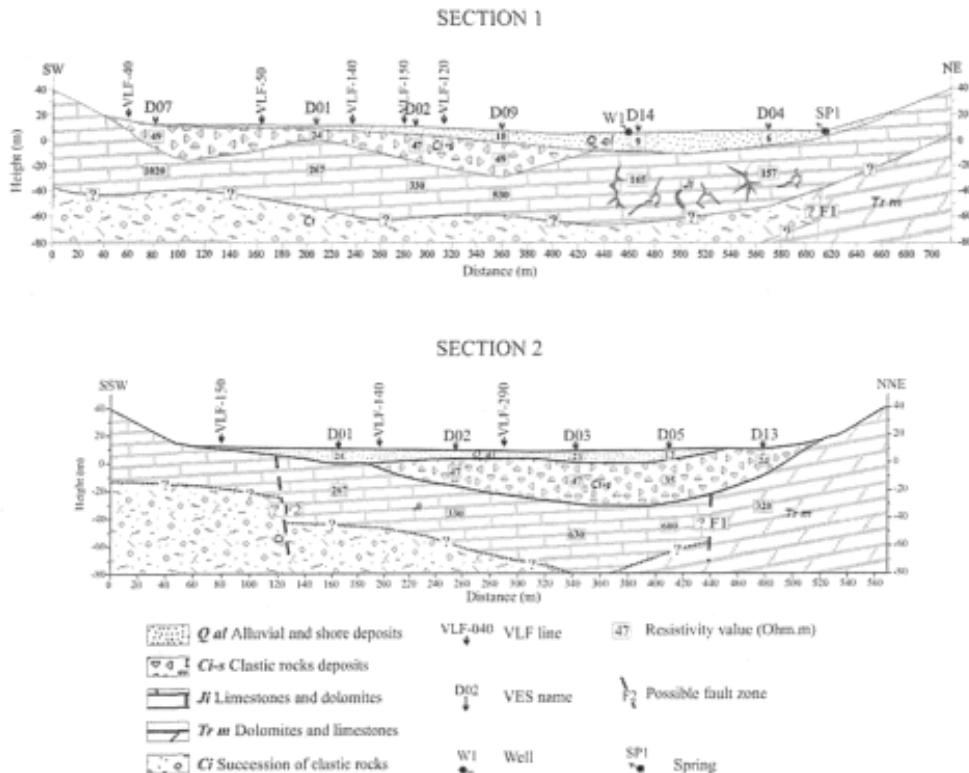


Figure 6. Integrated geological and geophysical data.

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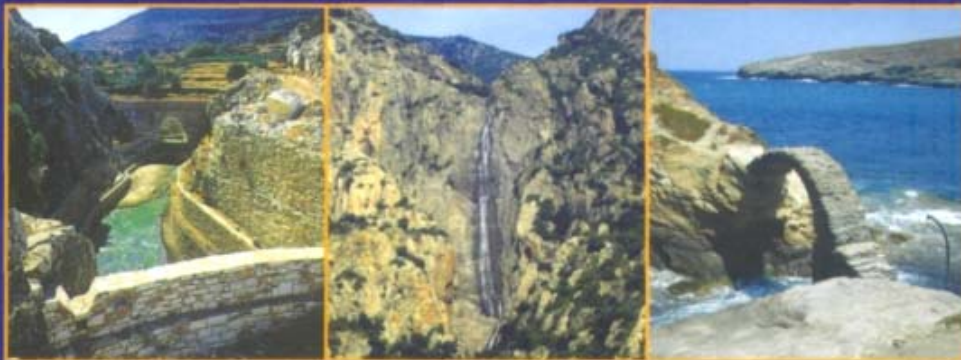
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**ΑΘΗΝΑ 2008 ATHENS**

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