

## PALEOBATHYMETRIC EVOLUTION OF THE EARLY LATE MIOCENE DEPOSITS OF THE PRE-APULIAN ZONE, LEVKAS ISLAND, IONIAN SEA

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### Abstract

*The Manassi section in Levkas Island belongs to the Pre-Apulian (Paxos) zone, the most external domain of the Hellenic realm. Its Early Tortonian sediments contain a rich foraminiferal fauna dominated, in numbers of individuals, by planktic species. Its benthic foraminiferal assemblage is characterized by a high number of taxa, with low numbers of individuals. Their study provides a basis for interpreting the paleobathymetry of the basin.*

*The Manassi section represents deposition in upper to lower bathyal depths, during a period of intense tectonic activity. Downslope transport of fauna by turbidity currents partly overprints the signal of paleobathymetrically-diagnostic foraminifera distribution. The recognition of allochthonous taxa is used, together with %P, to identify turbidite beds intercalated with in-situ marly sediments.*

*The micropaleontological and paleobathymetrical analyses of the studied sediments indicate that these correspond to distal atypical flysch deposited in the foredeep depozone of the most external domain (Pre-Apulian zone) of the Hellenide foreland basin.*

**Key words:** benthic foraminifera, paleobathymetry, Tortonian, Pre-Apulian zone, Mediterranean.

### Περίληψη

*Η τομή Μανάση, (νήσος Λευκάδα) ανήκει στην Προ-Απούλια (Παξοί) ζώνη, η οποία αποτελεί το εξωτερικότερο τμήμα των Ελληνίδων. Τα ιζήματα της τομής, ηλικίας Κατώτερο Τορτόνιο, περιέχουν μια πλούσια πανίδα τρηματοφόρων, όπου επικρατούν σε αριθμό ατόμων, τα πλαγκτονικά είδη. Η υψηλής ποικιλότητας συνάθροιση βενθονικών τρηματοφόρων χαρακτηρίζεται από έναν υψηλό αριθμό ειδών, με μικρό αριθμό ατόμων. Η ανάλυση της κατανομής των ειδών των βενθονικών τρηματοφόρων παρέχει στοιχεία για την παλαιοβαθυμετρική εξέλιξη της λεκάνης. Τα ιζήματα της τομής Μανάση αντιπροσωπεύουν απόθεση σε βάθη της ανώτερης έως και κατώτερης βαθύαλης ζώνης, η οποία έλαβε χώρα κατά τη διάρκεια της μετά του Μέσου*

Μειοκαίνου τεκτονικής ανύψωσης της περιοχής. Η κατά μήκος κατωφέρειας μεταφορά της πανίδας εξαιτίας της δραστηριότητας τουρβιδιτικών ρευμάτων επισκιάζει εν μέρει την κατανομή των διαγνωστικών για το παλαιοβάθος βενθονικών ειδών. Η αναγνώριση αλλόχθονων ειδών βενθονικών τρηματοφόρων χρησιμοποιείται, μαζί με την αναλογία των πλαγκτονικών τρηματοφόρων, για να καταδείξει την παρουσία τουρβιδιτικών στρωμάτων μέσα στις μάργες.

Η μικροπαλαιοντολογική και παλαιοβαθμετρική ανάλυση των αποθέσεων της τομής Μανάσση, έδειξε ότι αυτές αντιπροσωπεύουν έναν μακρινό άτυπο φλύσχη που αποτέθηκε στο κέντρο απόθεσης (*depo-center*) του πλέον εξωτερικού χώρου (Προ-Απούλιας ζώνη) της προχώρας των Ελληνίδων.

**Λέξεις κλειδιά:** βενθονικά τρηματοφόρα, παλαιοβαθμετρία, Τορτόνιο, Προ-Απούλια ζώνη, Μεσόγειος.

## 1. Introduction

A common goal in basin analysis is the assessment of paleobathymetry for use to interpret depositional environments, subsidence history and other aspects of basin evolution. The Middle Miocene foreland basin of Levkas island, Ionian Sea, needs considerable study both for its well-exposed outcrops of deep marine turbiditic strata and as an example of its sedimentary infill. Such studies are considerably hampered, however, by the lack of knowledge of paleobathymetry. The magnitude of uncertainty in paleobathymetric estimates could prevent a correct interpretation of subsidence analysis in such deep-water deposits.

Micropaleontologists have developed several methods to estimate paleodepths, in addition to methods based on the comparative abundance of planktic and benthic foraminifera. For example Culver (1988) documented the bathymetric distribution of some benthic foraminifera in the modern-day Gulf of Mexico, and used these to estimate paleodepths in some Cenozoic sequences in the same region. Murray and Alve (2000) found an overall trend towards an increase in benthic foraminiferal biodiversity with paleodepth. Van der Zwaan *et al.* (1990) determined a regression for the relationship between bathymetry and the percentage of planktic foraminifera with respect to the total fossil foraminiferal population (%P), based on present day bathymetric transects:

$$Depth(m) = e^{3.58718+(0.03534*\%P)}$$

where the plankton fraction  $\%P=100*P/(P+B-S)$ , and P is the amount of planktic foraminifera, B is the amount of benthic foraminifera and S is the amount of environmental stress markers, e.g. species of benthic foraminifera that can tolerate low oxygen concentrations [original list of Van der Zwaan *et al.* (1990) modified by Van Hinsbergen *et al.* (2005)].

However, the most widely applied methodology for assessing paleobathymetry is the analysis of benthic foraminiferal biofacies. In fact, benthic foraminifera have proved to be useful bathymetric indicators. Recent studies on modern faunas demonstrated that the faunal composition is strongly related to organic matter flux rates and dissolved oxygen concentrations in the bottom and pore water (e.g. Corliss 1985, Gooday 1994, Jorissen *et al.* 1995). In oceanic environments both parameters exhibit strong depth-dependent gradients. Consequently, certain benthic foraminiferal faunas are restricted to distinct bathymetric intervals. In particular, continental margins are characterized by a stenobathic distribution of many benthic foraminifera (e.g. Schmiedl *et al.* 1997).

In the present study, the overall nature of the benthic foraminiferal assemblages is examined to prove the validity of paleodepth estimations obtained computed using the percentage of planktic foraminifera, in part of the Miocene age deep sea sediments of Levkas Island.

## 2. Geological Setting

Levkas Island belongs to the Ionian Islands which are located in the west segment of the Hellenic Arc, the most active plate margin of the Mediterranean region (e.g. McKenzie 1978). The tectonic setting of the wider area is determined by the continental collision between northwestern Greece in the east and the Apulian platform in the west, as well as by the subduction of the African plate under the Aegean microplate along the active Hellenic Arc in the southwest. The Ionian Islands are situated in a transitional zone between the northwestern end of this active subsidence and the continental collision in the north.

The main tectonic structure of this transitional area is the Kephallinia Fault Zone (KFZ), which represents the active boundary between the SW-moving Aegean microplate and the Apulian platform. Based on data from Durmeijer *et al.* (1999) and Gautier *et al.* (1999) the age of the onset of the activity of KFZ could be estimated to a possible maximum age of the Pliocene-Pleistocene boundary. Many neotectonic active faults, striking to the NNE-SSW or E-W direction cross-cut the island (e.g. Cushing 1985).

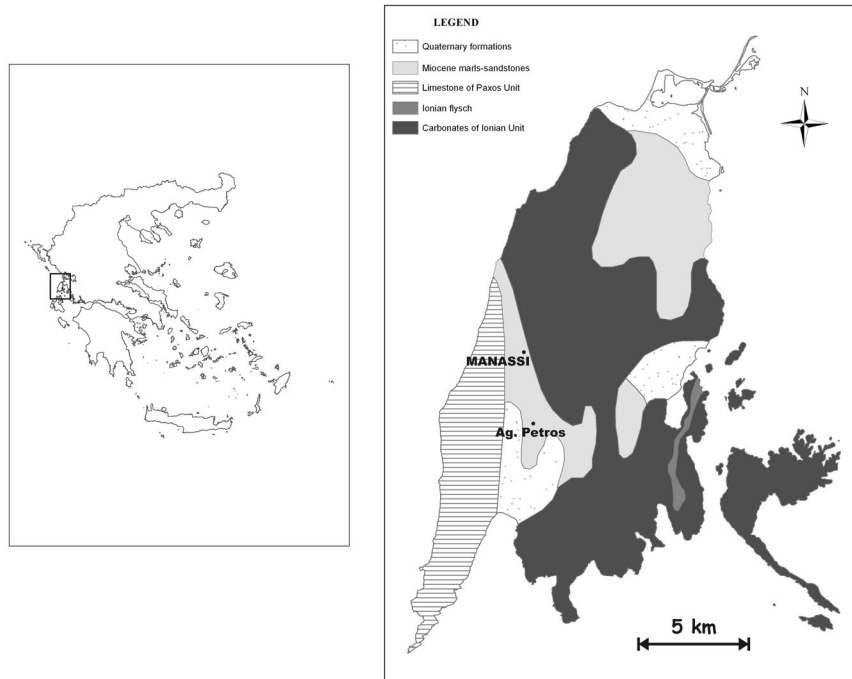
Levkas island is built up mainly by Alpine Mesozoic-Cenozoic sedimentary rocks belonging to the external units of the Hellenides, the Pre-Apulian (Paxos) zone and the Ionian zone (Fig. 1; Bornovas 1964, IGRS-IFP 1966, Cushing 1985), separated from one another by a major west-directed thrust fault (Aubouin 1957, Jacobshagen 1986). This thrust boundary is marked by evaporite intrusion suggesting that contractional deformation was the most important structural control of orogenesis in Western Greece (Karakitsios and Rigakis in press).

The Pre-Apulian zone corresponds to the most external domain of the Hellenic realm. It has been traditionally considered as a relatively uniform Mesozoic – Cenozoic carbonate transitional domain, between the Apulian platform and the Ionian basin (BP 1971, Karakitsios 1995). The general setting is complicated by strong tectonic deformation, including crustal extension, collision, and flexural subsidence, with undetermined amounts of shortening and block rotation (Karakitsios and Rigakis in press). Indeed, the outcropping successions differ in stratigraphic completeness, sedimentary development and faunal/floral content.

The depositional sequence in the Pre-Apulian zone begins with Triassic to Bajocian limestones, containing intercalations of black shales and anhydrite beds (borehole data, ESSO Hel. 1960). The lowest stratigraphic outcrops, located in Levkas Island, comprise Lower Jurassic dolomites and Middle Jurassic cherts and bituminous shales (Bornovas 1964, BP 1971). The Upper Jurassic consists of white chalky limestones and dolomite intercalations, with rare cherts and organic-carbon-rich black shales, containing planktic foraminifera. Lower Cretaceous limestones and dolomites outcrop only in Kephallonia Island, and are considered less pelagic in comparison to the Ionian facies of the same age. During the Campanian-Maastrichtian the platy limestones gradually become chalky with thin layers of argillaceous schists. They contain planktic foraminifera together with Rudists' fragments indicating the presence of intra-platform basins, characterizing the slope between the Apulian platform and the Ionian basin (Nikolaou 1986, Karakitsios and Rigakis in press).

The Paleocene formations consist of micritic limestones with planktic foraminifera (BP 1971). Intense tectonic activity, resulting in the differentiation of the Pre-Apulian zone, into relatively deep and shallow (sometimes emerged) areas, is deduced by sporadic hiatuses in between Paleocene and Late Cretaceous, described by Mirkou (1974). The Lower Eocene sequence comprises pelagic limestones with marl intercalations, whilst the Upper Eocene formations consist of reefal, unbedded limestones, with Algae, Bryozoans, Corals, Echinoids, and large foraminifera. Oligocene sediments were deposited in small basins (tectonic grabens) in between emerged areas, which were locally eroded. This setting can be attributed to the intense tectonic instability, which persisted throughout the Oligocene. During the Oligocene-Aquitainian, the intense diversification of foraminiferal assemblages (Accordi *et al.* 1998) suggests the presence of flexural subsiding

foreland basins (Karakitsios and Rigakis in press). In the late Early Miocene, progressive deepening ensued, with the demise of the former carbonate slope.



**Figure 1 - Simplified geological map of Levkas island indicating the location of the studied section (after Rondoyianni-Tsiambaou 1997)**

### 3. Study Location and Methods

This research deals with a more detailed study on paleobathymetry of a Miocene section (Manassi section) which belongs to the pre-Apulian zone and is located in the southwestern part of Levkas island (Fig. 1).

The Oligocene-Miocene deposits in this area conformably overlie Eocene limestones in the west (de Mulder 1975). Finely-bedded, pelagic limestones of Eocene age are overlain by a relatively thin succession of detrital, neritic limestones, passing upwards into at least 400-500 m of marls and clays, which may reach as high as the Upper Miocene (Bizon 1967).

The Manassi Section is located on the eastern slope of a N-S oriented valley. The studied succession consists of blue grey marls and clays with some fine grained sandstone interbeds. The intercalations of these thin, clastic beds and especially of positively graded sandstones in the studied succession reflect the influence of density currents, which supplied coarser material from a distant hinterland (de Mulder 1975).

The 25 m-thick section was measured and sampled at 0.5 to 1 m intervals. For the faunal analysis 26 samples (391-1 to 391-26) were taken (Fig. 2). Samples were washed over a 63  $\mu\text{m}$  sieve and subsequently the coarse fraction was dry sieved over 125  $\mu\text{m}$  sieve. The fraction  $>125 \mu\text{m}$  was then investigated for its benthic foraminiferal content. The picked planktonic specimens were counted and their (%P) was computed. Biostratigraphic attribution is based on the standard biozonal schemes of Iaccarino (1985) according to the absolute ages from Hilgen *et al.* (2000) and Foresi *et al.* (2002).

The total number of planktic (PFN) and benthic (BFN) foraminifera per gram was calculated by extrapolating the number of specimens in the split to the whole residue.

In order to calculate the changing paleobathymetry the following factors were identified:

Diversities were determined following the Shannon-Wiener index (e.g. Buzas and Gibson 1969). In addition, ratios between planktic (P) and benthic (B) foraminifera were calculated applying the formula  $P/(P+B)$ . The P/B ratios were integrated with marker benthic foraminifera in order to determine paleobathymetry.

To assess the depositional depth of the studied marine sediments, the general relationship between the fraction of planktic foraminifera with respect to the total foraminiferal population (%P) and depth of Van der Zwaan *et al.* (1990) was used, following sample selection and counting procedures described by van Hinsbergen *et al.* (2005).

The following depth zonation (McDougall, 1985) was used: inner shelf=0-50 m, outer shelf=50-150 m, upper bathyal=150-500 m, upper middle bathyal=500-1500 m, lower middle bathyal=1500-2000 m, lower bathyal=2000-3500 m, >3500 m=abyssal.

#### 4. Biostratigraphy

The planktonic foraminiferal analysis was carried out on the total of samples. Preservation was generally good, although at times was poor. The biostratigraphic results were based on the qualitative analysis of planktonic foraminifera, analyzed at the species level. Their specific identifications were conducted according to reference publications (Dermitzakis 1978, Iaccarino 1985, Hilgen *et al.* 2000, Foresi *et al.* 2002). For the biostratigraphic analysis we only considered the appearance/disappearance of marker species and paracme/acme intervals of selected taxa according to Foresi *et al.* (2002).

The qualitative analysis revealed 14 planktonic foraminifera categories: *Paragloborotalia siakensis*, *Globigerinoides obliquus obliquus*, *Globorotalia menardii-preamenardii*, *Catapsydrax parvulus*, *Neogloboquadrina atlantica praeatlantica*, *P. partimlabiata* and *N. acostaensis*. The coiling direction of the neogloboquadriniids was investigated separately because of their biostratigraphic value.

*G. obliquus obliquus* is present in all the samples and no specimens of *G. subquadratus* were observed. The last common occurrence (LCO) of *G. subquadratus*, has been dated in the Mediterranean at 11.54 and 11.58 Ma (Hilgen *et al.* 2000, 2003, Lirer *et al.* 2002) and at same time (11.54 Ma according to Hilgen *et al.* 2000, 2003, Lirer *et al.* 2002), the first regular occurrence (FRO) of *G. obliquus* is recorded. This event coincides with the end of the first influx of the neogloboquadrinids (FO of neogloboquadrinids 11.78 Ma), (Hilgen *et al.* 2000, 2003, Foresi *et al.* 2002). Above this level *N. acostaensis* occurs in low percentages.

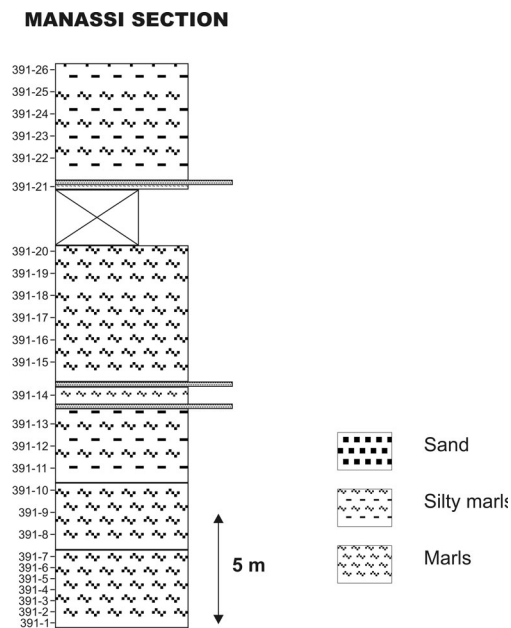


Figure 2 - Lithostratigraphical column of the studied section

In our record few specimens of *N. acostaensis* are recorded (2-5 %) with random (left and right) coiling direction. *N. atlantica praeatlantica* is also present having the same distributional pattern of *N. acostaensis*. A paracme interval in the distribution range of neogloboquadriniids has been recorded in Mediterranean and dated at 11.54 -11.21 Ma (Foresi *et al.* 2002). These dates coincide with the LCO of *G. subquadratus* and the LO of *P. siakensis* respectively (Iaccarino *et al.* 2004, Foresi *et al.* 2002, Hilgen *et al.* 2003). In the same interval *C. parvulus* is abundant (Foresi *et al.*, 2002). In the studied section the same pattern of *C. parvulus* is observed.

*P. siakensis* left coiled is abundant from the base of the section upward. The last occurrence of this species has been dated at 11.21 Ma (Hilgen *et al.* 2000, Caruso *et al.* 2002, Di Stefano *et al.* 2002, Foresi *et al.* 2002, Lirer *et al.* 2004, Iaccarino *et al.* 2004). *Catapsydarx parvulus* is present in several samples throughout the section. The last occurrence of this species in Mediterranean is reported at 9.91 Ma (Hilgen *et al.* 2000).

Consequently, according to the biostratigraphical criteria such as the distributional pattern of neogloboquadriniids, the regular occurrence of *G. obliquus obliquus* and the presence of *P. siakensis* in all the samples, the section covers the time interval between 11.54 and 11.2 Ma, having a lower Tortonian age, just above the Serravallian/Tortonian boundary astronomically dated at 11.6 Ma (Hilgen *et al.* 2005).

## 5. Benthic Foraminifera Species

Of the 26 samples, only one was barren of benthic foraminifera. It is sample 391-26 which contains high fractions of quartz and rock fragments, which probably derived from downslope transport. Four samples yielded fewer than 200 benthics (391-21, 105 total foraminifera only, 391-23, 180 foraminifera only, 391-24, 183 foraminifera only and 391-25, 73 foraminifera only). However, paleodepths for these samples were estimated.

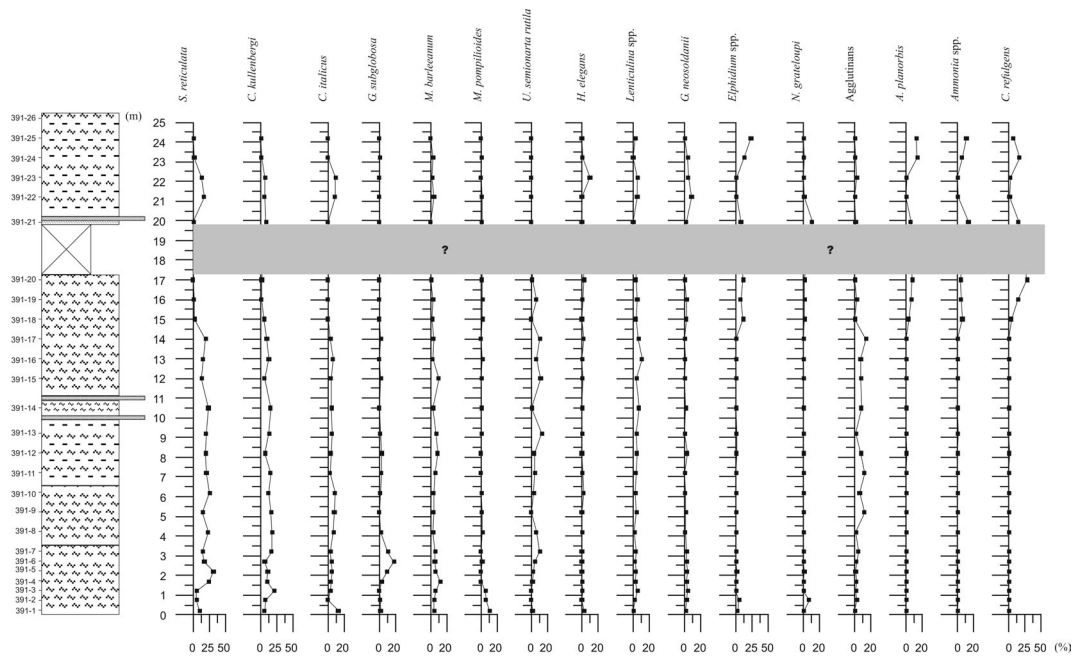
One hundred and sixty nine species and species groups of benthic foraminifera were identified in the 25 samples studied. The primary taxonomic references are given in Ellis and Messina (1940 and supplements).

Fig. 3 shows the frequency of the dominant benthic species in the samples. The most abundant species recovered from the section were *Siphonina reticulata*, *Cibicidoides kullenbergi*, *Melonis barleeanus*, *C. italicus*. These taxa abundantly and persistently occur throughout the section along with *Globocassidulina subglobosa*, *Gyroidinoides neosoldanii* and *Uvigerina semiornata rutila*. In the upper part of the section, clearly transported shallow water taxa co-occur with deep water species. In these cases it can be clearly inferred that the shallow water taxa have been dislodged downslope into deeper waters, through transport or turbidity activity.

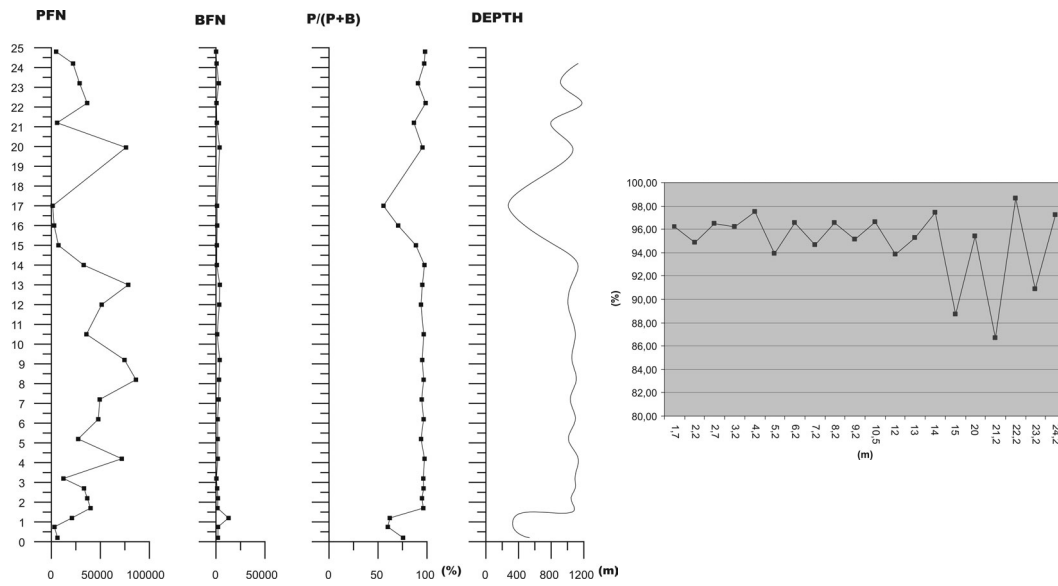
## 6. Paleobathymetry

### 6.1. Analysis of %P planktic foraminifera

The result of total planktic foraminiferal (PFN) fauna performed in this study is shown in Fig.4, together with BFN trends. With the exception of samples 391-1 to 391-3 and 391-20, the planktic abundance (P/P+B) in all observed Manassi samples is greater than 90 %. According to Van der Zwaan *et al.* (1990) these data suggest an average water depth of about 1000 m. Moreover, the PFN strongly fluctuates between 1330 specimens per gram and 78.500 specimens per gram; whereas BFN remains relatively stable exhibiting low values (Fig. 4). According to Mendes *et al.* (2004), the BFN is inversely related with water depth. Slightly increased BFN in the lowermost part of the section (sample 391-3) indicates that at this stratigraphic level bottom water conditions have been generally more favorable for the production of benthic foraminifera.



**Figure 3 - Abundance pattern of the most significant benthic foraminifera species identified in the studied section**



**Figure 4 - PFN, BFN, P/B ratios together with the paleodepth curve obtained after the application of the regression of Van der Zwaan *et al.* (1990) of the studied samples. The inlet to the right represents variations in planktic percent above 90 exclusive three samples which show values below 60 %**

Fig. 4 also shows planktic percentages of Manassi section samples exclusive of the four aforementioned samples. Several fluctuations are revealed by this graph. There is no relationship between these fluctuations and the sediment grain size as all the samples are from the grey marls. Although the significance of variations in planktic percent above 90 % is problematical, this may

merely reflect periodic dilution of the autochthonous deep-water fauna with transported shallower faunas (e.g. Robertson 1998). It is likely that, in these beds, some degree of faunal reworking has occurred, displacing shallow water taxa into deeper water faunas. This condition suggests that these beds could correspond to turbidite horizons.

In samples 391-1 to 391-3 and 391-20 there is a drop in P/B together with a drop in PFN. According to Speijer and Schmitz (1998), where fluctuations between these records coincide, it is likely that carbonate dissolution has occurred at some step during the generation of the fossil assemblage. In exceptional situations this association of changes may indicate a real paleoenvironmental change. If so, in shelf settings this should be readily discernable from the composition of the planktic and/or benthic foraminiferal assemblages.

Application of van der Zwaan *et al.* (1990) regression indicates that deposition began at ~530 m (deep upper bathyal) during the deposition of sample 391-1 (0.2 m) and regressed to ~340 m (upper bathyal) in sample 391-3 (1.2 m). The sea then transgressed to a paleodepth of ~1130 m (deep lower middle bathyal) during the deposition of sample 391-17 (14 m). An ensuing rapid regression saw sea-levels attain a paleodepth ~270 m (upper bathyal) during the deposition of sample 391-20, then transgress abruptly to ~1100 m (deep middle bathyal) at the time of sample 391-21 and remained relatively constant with short-term fluctuations from sample 391-22 till 391-25.

## 6.2. Diversity

Shannon-Wiener H(S) values in the Manassi section vary from 2.61 to 3.42 (Table 1). Gibson and Buzas (1973) showed that these values correlate to water depths ranging from 80 to 3200 m in the northeastern Gulf of Mexico. In view of the wide ranges of water depths associated with diversity measure H(S), it is unrealistic to assign absolute water depths to Manassi deposition based on these criteria, although increasing diversity probably indicates increasing water depth in a relative sense. Therefore the variations in diversity may reflect only variations in degree of dilution by shallower water forms.

## 6.3. Bathymetric significance of foraminiferal assemblages

Any significant trends in palaeobathymetry can be independently checked on the studied samples along the succession, by identifying marker species for selected depth intervals.

The bathymetric preferences of the different foraminiferal taxa were assessed by evaluating their distribution pattern in modern oceans (Table 1). Based on this information the bathymetric evolution at the investigated site can be reconstructed. The paleodepths indicated by the benthic foraminifera in samples 391-4 to 391-18, 391-22 and 391-23 agree with those calculated using Van der Zwaan *et al.* (1990) regression. The benthic assemblages in these samples indicate an upper middle bathyal environment, being dominated by *C. kullenbergi* and *S. reticulata*. For samples 391-18 to 391-21, 391-24 and 391-25, however, the paleodepths calculated from the percentage of planktic foraminiferal fauna differ markedly from those indicated by the benthic one. These samples yielded benthic assemblages that were dominated by *Elphidium* spp. and *C. refulgens* associated with a lesser *A. planorbis* and *Ammonia* spp. The normally shallow-water preferences of *Elphidium* species (Hayward *et al.* 1997) and *C. refulgens* (Murray 1991) suggest that these samples were deposited at neritic paleodepths and not at the bathyal paleodepths as suggested by the abundant planktic foraminifera. Furthermore, benthic foraminifera typical of deep-water habitats, such as were recorded in the lower part of the sequence, were also recovered from these samples, yet in small percentages. Thus the presence of specimens from *Elphidium* species and *C. refulgens* in fossil bathyal assemblages should be explained by sediment displacement (e.g. Phleger *et al.* 1953) or rafting of plant material to which cibicidids lived attached into pelagic environments after storms (Sprovieri and Hasegawa 1990).



**Table 1- Depth distribution of a selection of benthic marker species based on Parker (1958), Blanc-Vernet (1969), Wright (1978), Parisi (1981), Jorissen (1987), Sprovieri and Hasegawa (1990), Sgarella and Moncharmont Zei (1993), De Stigter *et al.* (1998), De Rijk *et al.* (2000), Seidenkrantz *et al.* (2000), Jannink (2001), Kouwenhoven *et al.* (2003). Note that these studies are all based on the Mediterranean area**

Samples	Predominant Species	Abundant Species	Diversity	%P	Depth	Depth (m)	
391-1	<i>C. italicus</i>	<i>M. pompilioides</i> <i>S. reticulata</i>	3,17	75,54	Deep Upper Bathyal	532,30	
391-2	<i>C. kullenbergi</i>	<i>A. helycinus</i> <i>M. pompilioides</i>	3,41	60,00	Upper Bathyal	332,23	
391-3	<i>C. kullenbergi</i>		3,07	62,05		342,38	
391-4	<i>S. reticulata</i>	<i>C. kullenbergi</i> <i>M. barleeanus</i>	2,88	96,22	Deep Lower Bathyal	1085,11	
391-5	<i>S. reticulata</i>	<i>C. kullenbergi</i> <i>G. subglobosa</i>	2,67	94,90		1043,03	
391-6	<i>G. subglobosa</i>	<i>S. reticulata</i> <i>C. kullenbergi</i>	2,97	96,49		1094,18	
391-7	<i>C. kullenbergi</i>	<i>S. reticulata</i> <i>G. subglobosa</i> <i>U. rutila</i>	2,83	96,21		1090,81	
391-8	<i>S. reticulata</i>	<i>C.kullenbergi</i>	2,75	97,50		1131,74	
391-9	<i>C. kullenbergi</i>	<i>S.reticulata</i>	3,03	93,91		1010,36	
391-10	<i>S. reticulata</i>	<i>C.kullenbergi</i>	2,68	96,54		1096,30	
391-11	<i>S. reticulata</i>	<i>C.kullenbergi</i>	3,1	94,65		1033,51	
391-12	<i>S. reticulata</i>	<i>C.kullenbergi</i>	3,12	96,57		1108,00	
391-13	<i>S. reticulata</i>	<i>C.kullenbergi</i> <i>U. rutila</i>	2,8	95,17		1053,70	
391-14	<i>S. reticulata</i>	<i>C.kullenbergi</i> <i>A. helycinus</i>	2,71	96,63		1095,82	
391-15	<i>S. reticulata</i>	<i>U. rutila</i>	3,32	93,84		1005,78	
391-16	<i>S. reticulata</i>	<i>C.kullenbergi</i> , <i>Lenticulina</i> spp.	2,99	95,24		1047,53	
391-17	<i>S. reticulata</i>	<i>C.kullenbergi</i> <i>U. rutila</i>	3,18	97,42		1128,66	
391-18	<i>Elphidium</i> spp.	<i>Ammonia</i> spp.	3,42	88,69		Upper Middle Bathyal	848,01
391-19	<i>C. refulgens</i>	<i>Elphidium</i> spp. <i>A. planorbis</i>	3,41	70,54			475,97
391-20	<i>C. refulgens</i>	<i>Elphidium</i> spp. <i>A. planorbis</i>	2,93	55,56		Upper Bathyal	274,82
391-21	<i>C. refulgens</i>	<i>Elphidium</i> spp.	2,79	95,39	Deep	1063,64	

		<i>Ammonia</i> spp.			Middle Bathyal	
391-22	<i>S. reticulata</i>	<i>C. kullenbergi</i> , <i>G. neosoldanii</i>	3,23	86,70		793,49
391-23	<i>S. reticulata</i>	<i>C. ungerianus</i> , <i>H. elegans</i>	3,05	98,62		1178,38
391-24	<i>C. refulgens</i>	<i>Elphidium</i> spp. <i>A. planorbis</i>	3,02	90,85		914,46
391-25	<i>C. refulgens</i>	<i>A. planorbis</i> , <i>Ammonia</i> spp.	2,61	97,20	1128,89	

## 7. Discussion-Conclusions

The present study was undertaken to provide a comprehensive paleobathymetric analysis of a section (Manassi section) located in the central part of the pre-Apulian Zone of Levkas Island. The biostratigraphic analysis based on planktic foraminifera indicated a Lower Tortonian age (between 11.54 and 11.21 Ma). This time interval is considered very crucial for the studied area as it marks the transition from carbonate to clastic sedimentation in western Greece and a phase of compression which affected the external Ionian zone and the pre-Apulian zone of Levkas. This shortening is related to the overriding of the pre-Apulian zone by the Ionian zone.

The paleobathymetric analysis showed that sedimentation occurred in a wide, deep and well-ventilated foreland basin. Deposition occurred on a steep, unstable slope adjacent to a narrow shelf. Periodic downslope transport of shallow-water sediments into deeper water occurred throughout this interval, perhaps related to intense tectonic activity. The prominent occurrence of turbidite horizons can be linked to flysch deposition.

It has so far been generally accepted, that the Pre-Apulian zone lacks typical flysch sediments. However, we consider the post - Oligocene deposits as the distal atypical flysch sediments, deposited in the foredeep depozone of the most external domain (Pre-Apulian zone) of the Hellenides' foreland basin system (DeCelles and Giles 1996). Their partial absence in some areas is due to the fact that these areas correspond to the most external foreland basin forebulge uplifts, which sediments have possibly been eroded, leaving only unconformities.

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